

Modeling azimuthal asymmetries of the troposphere delay during a 14-days typhoon period in Tsukuba

A. Pany, J. Boehm, T. Hobiger, R. Ichikawa, H. Schuh

19th EVGA Working Meeting, Bordeaux, France
March 24-25, 2009

Motivation

- state-of-the-art modeling of troposphere delay:

$$\Delta L = \Delta L_h^z \cdot mf_h(e) + \Delta L_w^z \cdot mf_w(e) + \\ mf_g(e) \cdot [G_N \cdot \cos(\alpha) + G_E \cdot \sin(\alpha)]$$

ΔL

troposphere total delay

Motivation

- state-of-the-art modeling of troposphere delay:

$$\Delta L = \Delta L_h^z \cdot mf_h(e) + \Delta L_w^z \cdot mf_w(e) + mf_g(e) \cdot [G_N \cdot \cos(\alpha) + G_E \cdot \sin(\alpha)]$$

ΔL

troposphere total delay

symmetric part

asymmetric part

Motivation

- state-of-the-art modeling of troposphere delay:

$$\Delta L = \Delta L_h^z \cdot mf_h(e) + \Delta L_w^z \cdot mf_w(e) + mf_g(e) \cdot [G_N \cdot \cos(\alpha) + G_E \cdot \sin(\alpha)]$$

ΔL

troposphere total delay

$\Delta L_h^z / \Delta L_h^w$

zenith hydrostatic/wet delay

mf_h / mf_w

mapping function hydrostatic/wet

e

elevation angle

mf_g

$mf_h \cdot \cot(e)$ or $mf_w \cdot \cot(e)$

G_N / G_E

north/east gradient

α

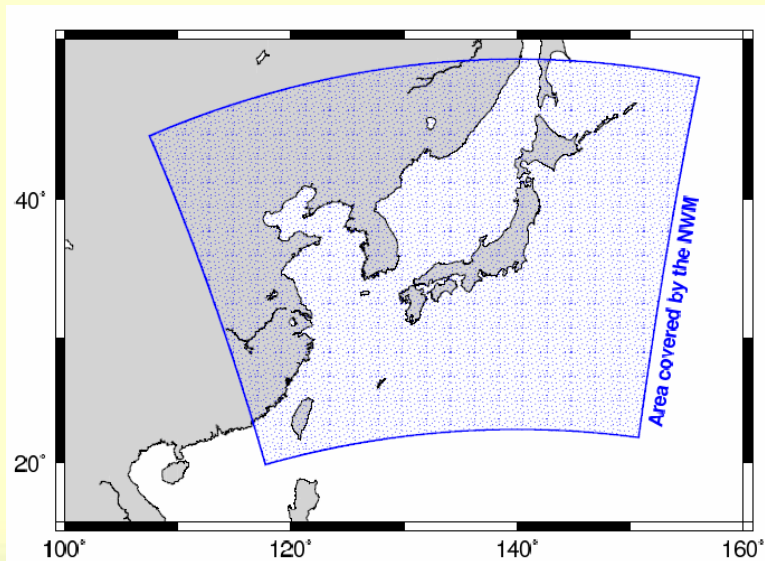
azimuth

The idea of this work

- high resolution numerical weather models (NWM) allow ray-tracing through atmosphere
- Japan Meteorological Agency (JMA) provides Meso-scale Analysis Data (MANAL)

The idea of this work

- high resolution numerical weather models (NWM) allow ray-tracing through atmosphere
- Japan Meteorological Agency (JMA) provides Meso-scale Analysis Data (MANAL)
- covers Japan, Korea, Taiwan, and Eastern China



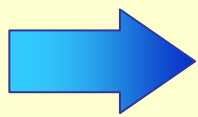
Real-time ray-tracing through numerical weather models for space geodesy, T. Hobiger, R. Ichikawa, Y. Koyama, T. Kondo, NICT IVS Technical Development Center News, 2008

The idea of this work

- high resolution numerical weather models (NWM) allow ray-tracing through atmosphere
- Japan Meteorological Agency (JMA) provides Meso-scale Analysis Data (MANAL)
- covers Japan, Korea, Taiwan, and Eastern China
- ray-traced delays can be used to correct geodetic observations

The idea of this work

- high resolution numerical weather models (NWM) allow ray-tracing through atmosphere
- Japan Meteorological Agency (JMA) provides Meso-scale Analysis Data (MANAL)
- covers Japan, Korea, Taiwan, and Eastern China
- ray-traced delays can be used to correct geodetic observations



use ray-traced delays to improve mapping functions and explore possibilities for improving the modeling of azimuthal asymmetries

Data

- ray-traced delays were computed with the **KA**shima **RA**y-tracing **T**ools (KARAT) (Hobiger et al., 2008)
- from meteorological fields of the **J**apan **M**eteorological **A**gency (JMA)

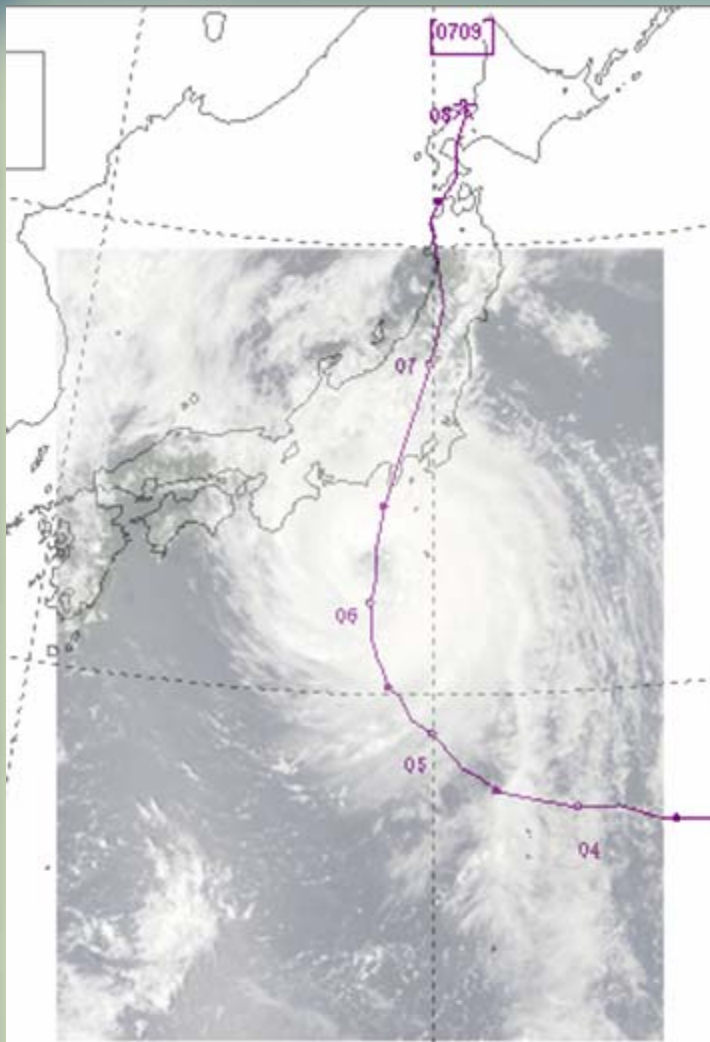
Data



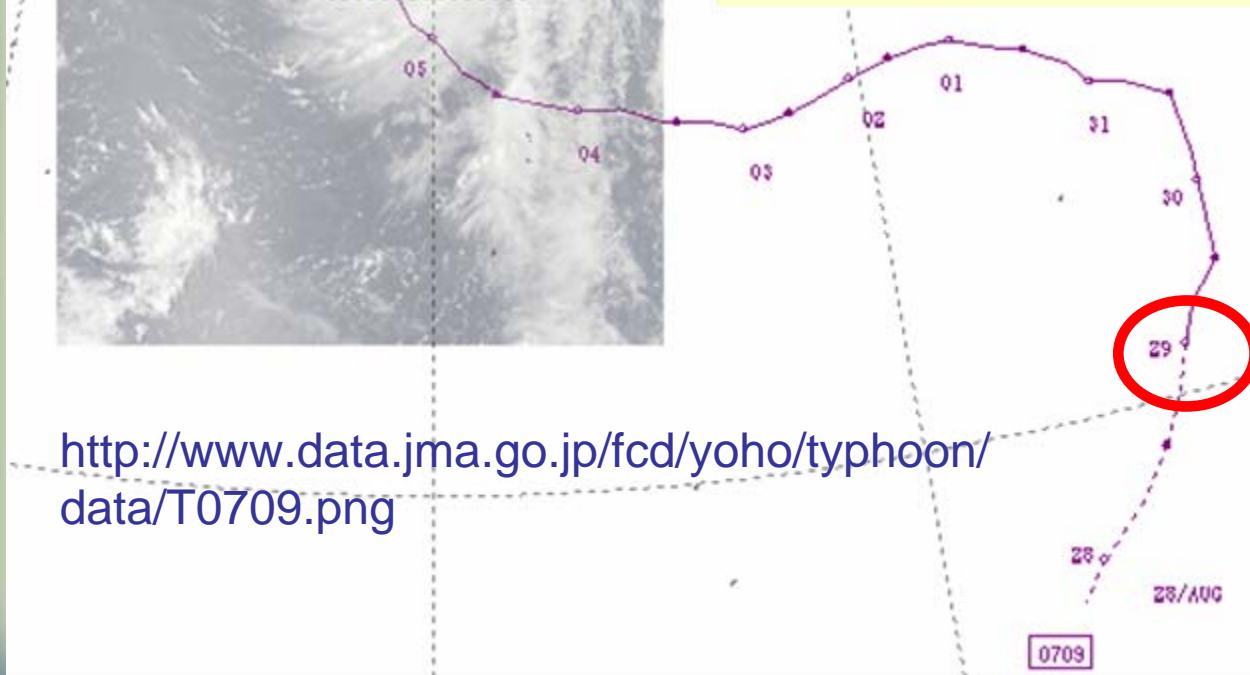
- typhoon „Fitow“

photo taken on 06/09/2007,
1:15 pm local time

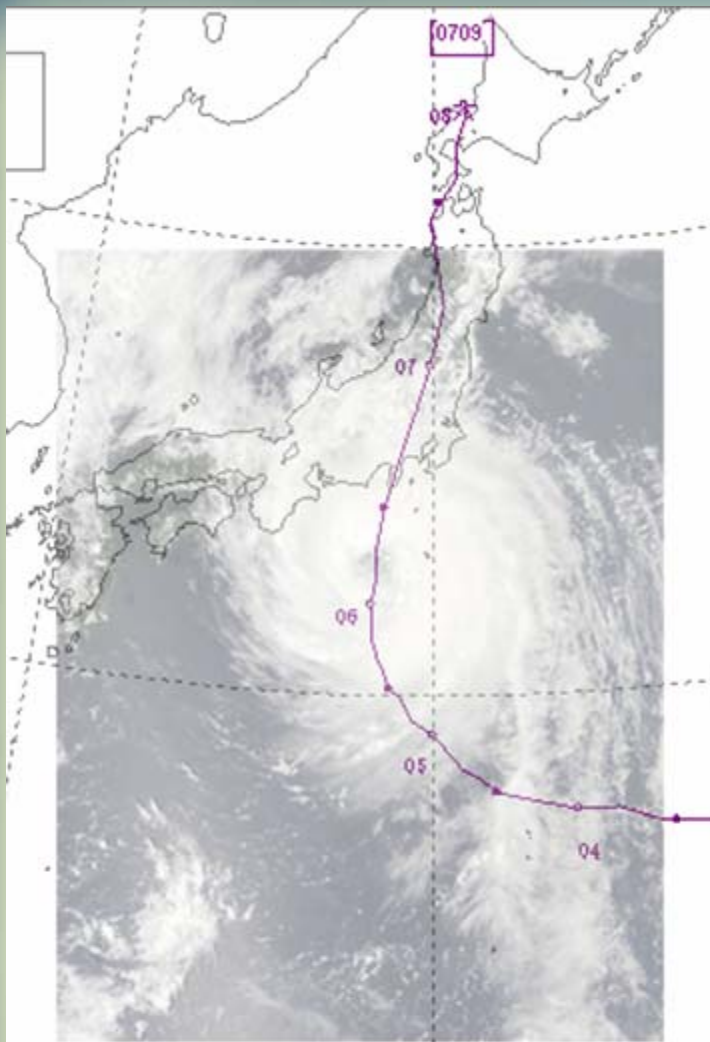
<http://earthobservatory.nasa.gov/NaturalHazards/view.php?id=18973>



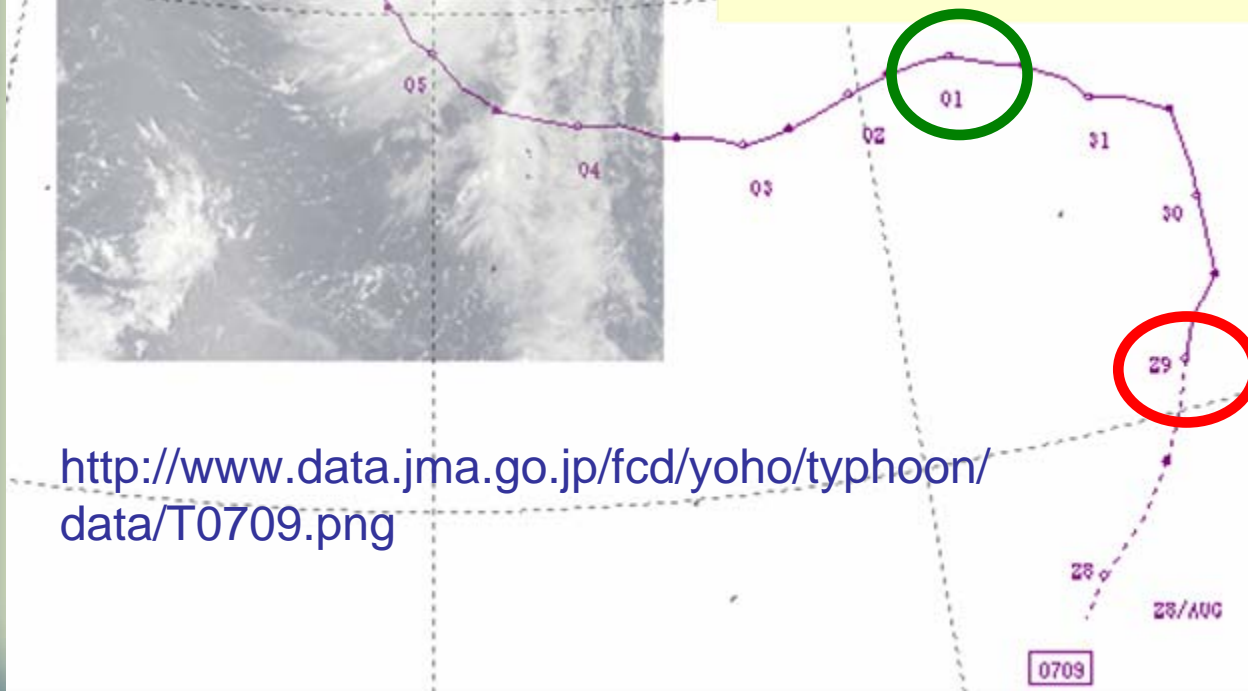
- typhoon „Fitow“
- named typhoon on August 29, 2007



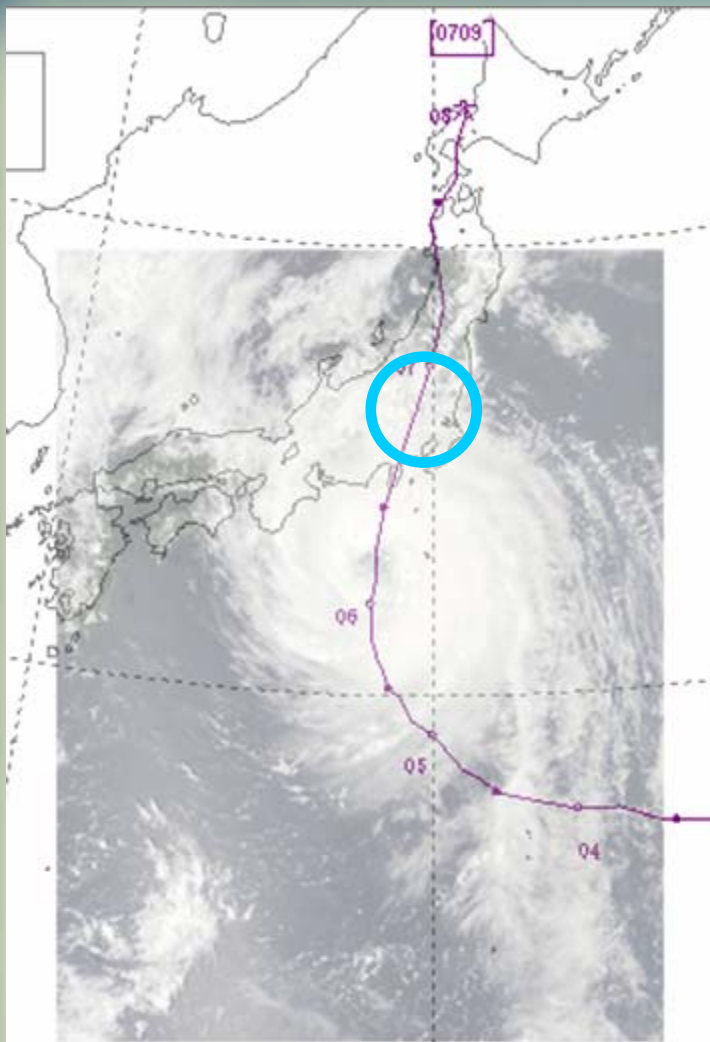
<http://www.data.jma.go.jp/fcd/yoho/typhoon/data/T0709.png>



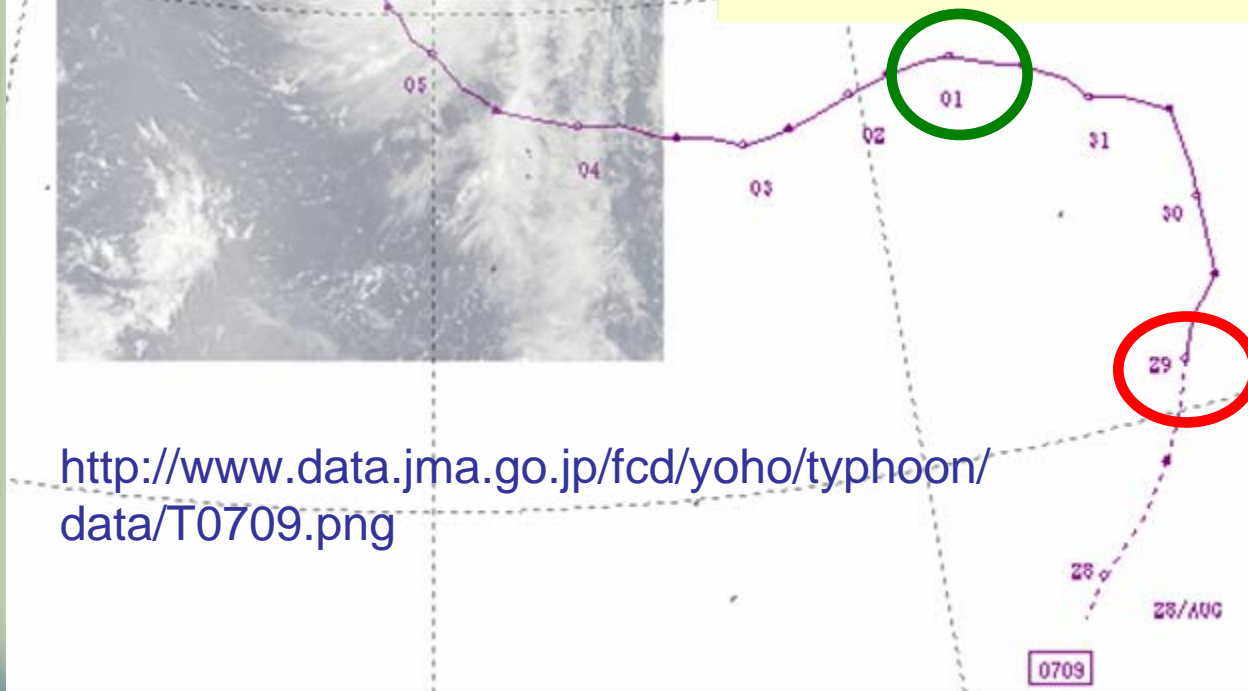
- typhoon „Fitow“
- named typhoon on August 29, 2007
- collected during typhoon period
- September 1 – September 14



<http://www.data.jma.go.jp/fcd/yoho/typhoon/data/T0709.png>



- typhoon „Fitow“
- named typhoon on August 29, 2007
- collected during typhoon period
- September 1 – September 14



passed near
Tsukuba on
September 7

<http://www.data.jma.go.jp/fcd/yoho/typhoon/data/T0709.png>

Data

- ray-traced delays were computed with the **KA**shima **RA**y-tracing **T**ools (KARAT) (Hobiger et al., 2008)
- from meteorological fields of the **J**apan **M**eteorological **A**gency (JMA)
- lowest outgoing elevation angle is 3° ($\sim 3.3^\circ$ on the ground)

Data

- ray-traced delays were computed with the **KA**shima **RA**y-tracing **T**ools (KARAT) (Hobiger et al., 2008)
- from meteorological fields of the **J**apan **M**eteorological **A**gency (JMA)
- lowest outgoing elevation angle is 3° ($\sim 3.3^\circ$ on the ground)
- 1° resolution in elevation and azimuth

Data

- ray-traced delays were computed with the **KA**shima **RA**y-tracing **T**ools (KARAT) (Hobiger et al., 2008)
- from meteorological fields of the **J**apan **M**eteorological **A**gency (JMA)
- lowest outgoing elevation angle is 3° ($\sim 3.3^\circ$ on the ground)
- 1° resolution in elevation and azimuth
- **3h temporal resolution**

Data

- ray-traced delays were computed with the **KA**shima **RA**y-tracing **T**ools (KARAT) (Hobiger et al., 2008)
- from meteorological fields of the **J**apan **M**eteorological **A**gency (JMA)
- lowest outgoing elevation angle is 3° ($\sim 3.3^\circ$ on the ground)
- 1° resolution in elevation and azimuth
- 3h temporal resolution
- delays exhibit significant azimuthally asymmetric characteristics

Mapping Function

continued fraction form

$$mf(e) = \frac{1 + \frac{a}{1 + \frac{b}{1 + c}}}{\sin(e) + \frac{a}{\sin(e) + \frac{b}{\sin(e) + c}}}$$

Mapping Function

continued fraction form

$$mf(e) = \frac{1 + \frac{a}{1 + \frac{b}{1 + c}}}{\sin(e) + \frac{a}{\sin(e) + \frac{b}{\sin(e) + c}}}$$

how coefficients are determined depends on the mapping function that is used

- e.g. from ray-traced delays

Vienna Mapping Function 1 (VMF1)

remove hydrostatic delay and estimate
wet mapping function

Vienna Mapping Function 1 (VMF1)

remove hydrostatic delay and estimate
wet mapping function



b and ***c*** from GMF, ***b*** = 0.0029,
c computed as function of day of year and latitude

Vienna Mapping Function 1 (VMF1)

remove hydrostatic delay and estimate
wet mapping function



b and ***c*** from GMF, ***b*** = 0.0029,
c computed as function of day of year and latitude



for each azimuth

fix ***a*** at 3 deg
elevation

determine ***a*** with adjust-
ment over all elevations

Vienna Mapping Function 1 (VMF1)

remove hydrostatic delay and estimate
wet mapping function



b and ***c*** from GMF, ***b*** = 0.0029,
c computed as function of day of year and latitude



for each azimuth

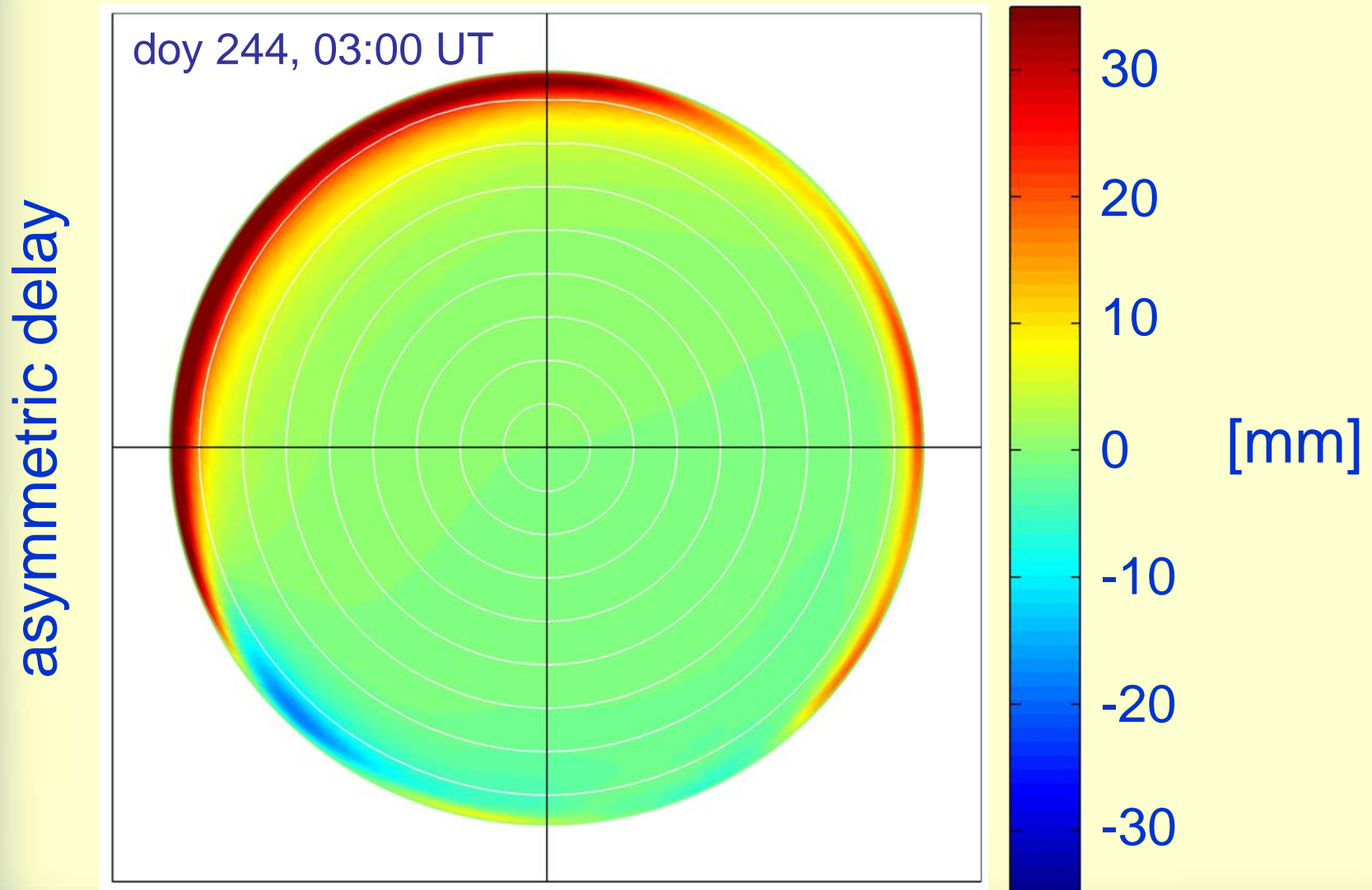
fix ***a*** at 3 deg
elevation

determine ***a*** with adjust-
ment over all elevations

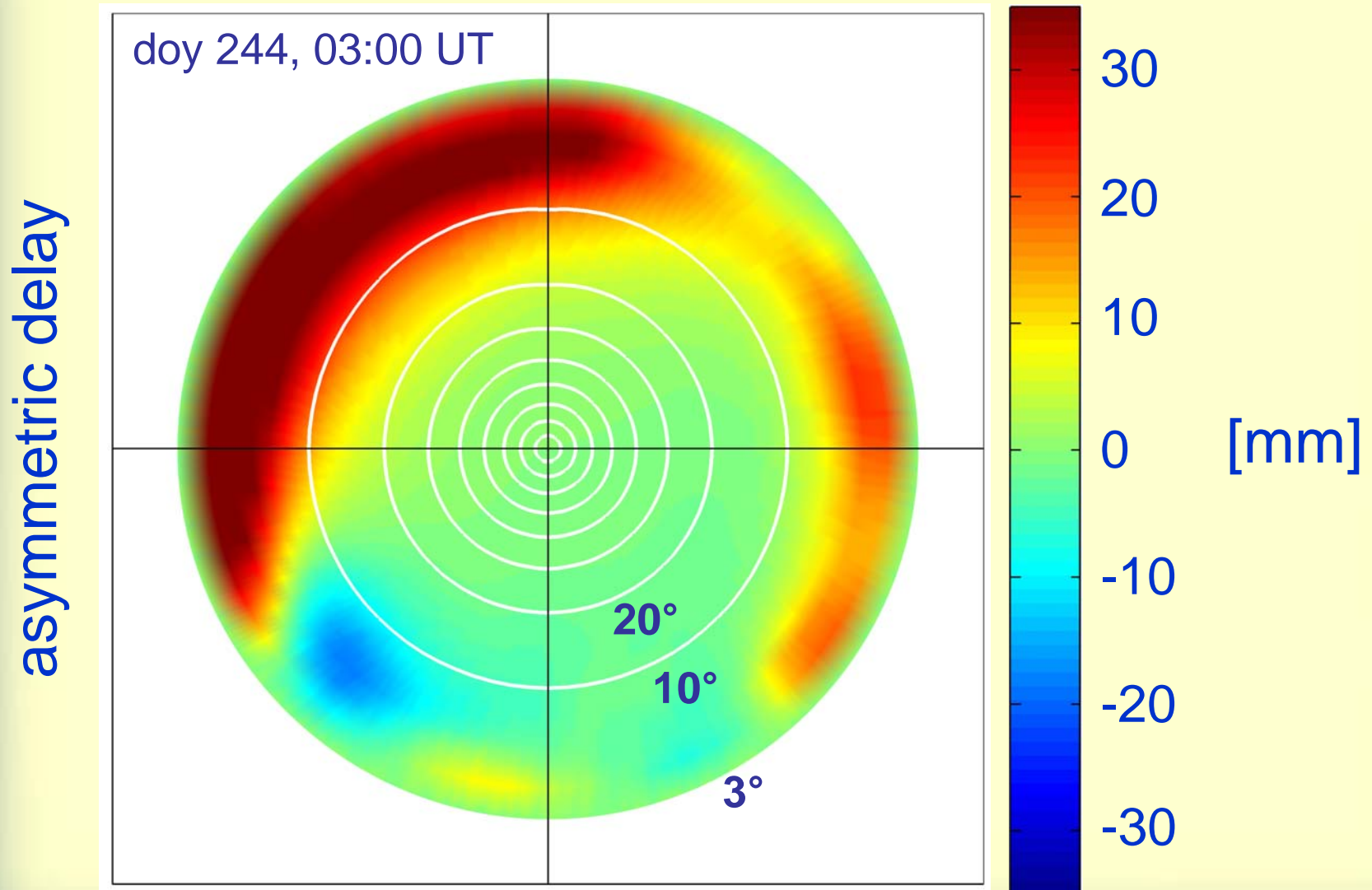


mean ***a*** (over all α)

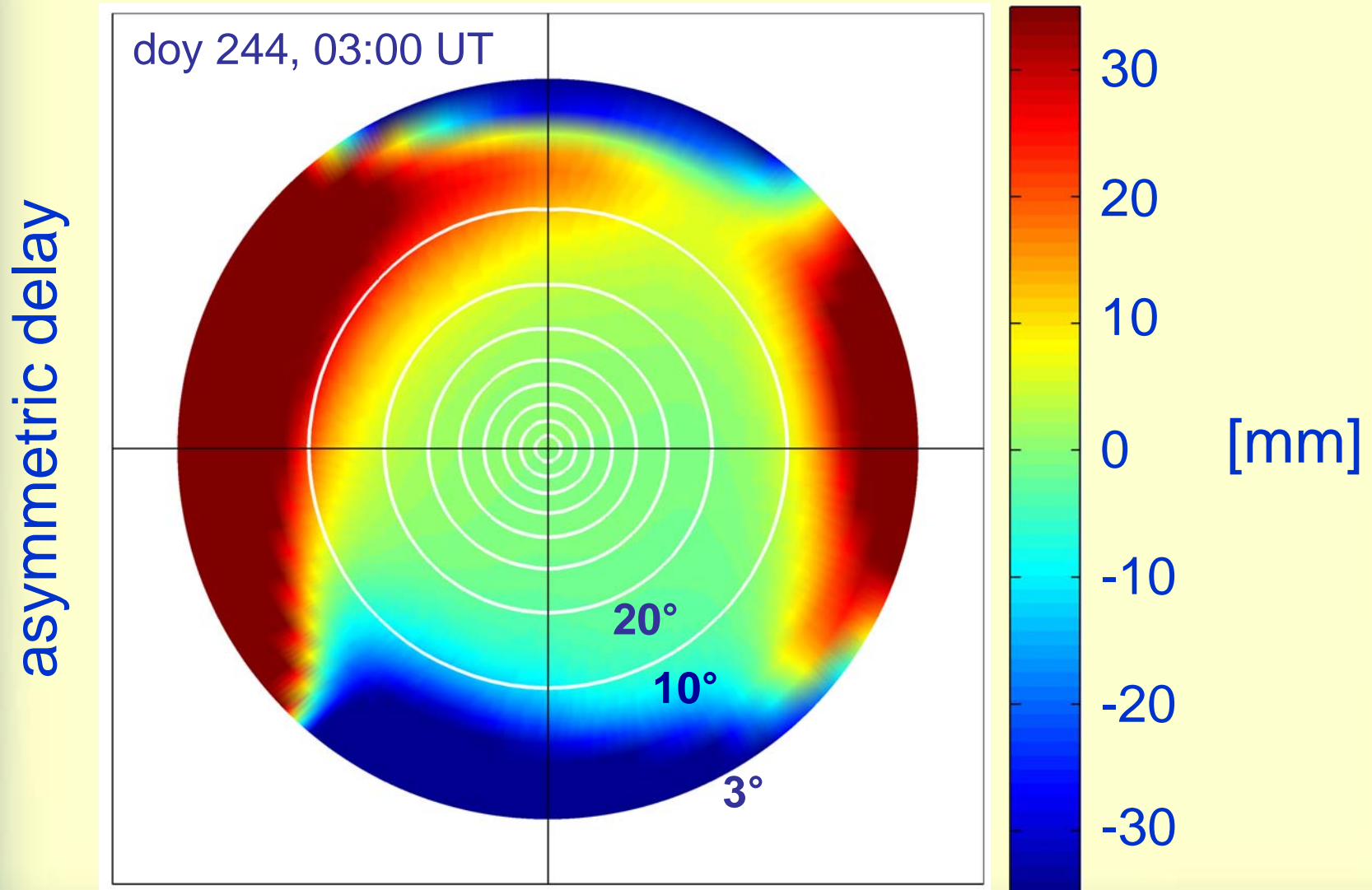
a at 3° for each azimuth



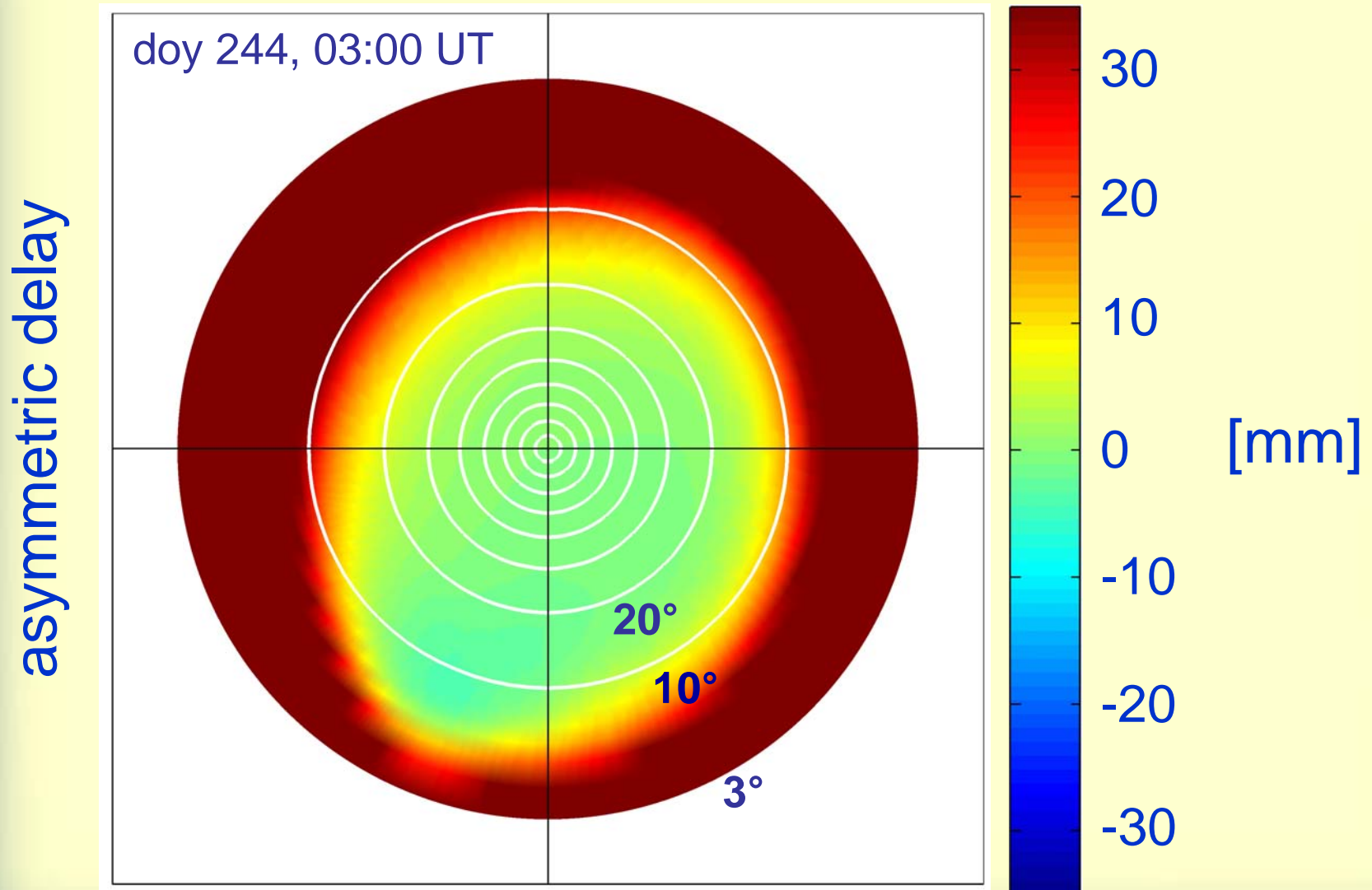
a at 3° for each azimuth



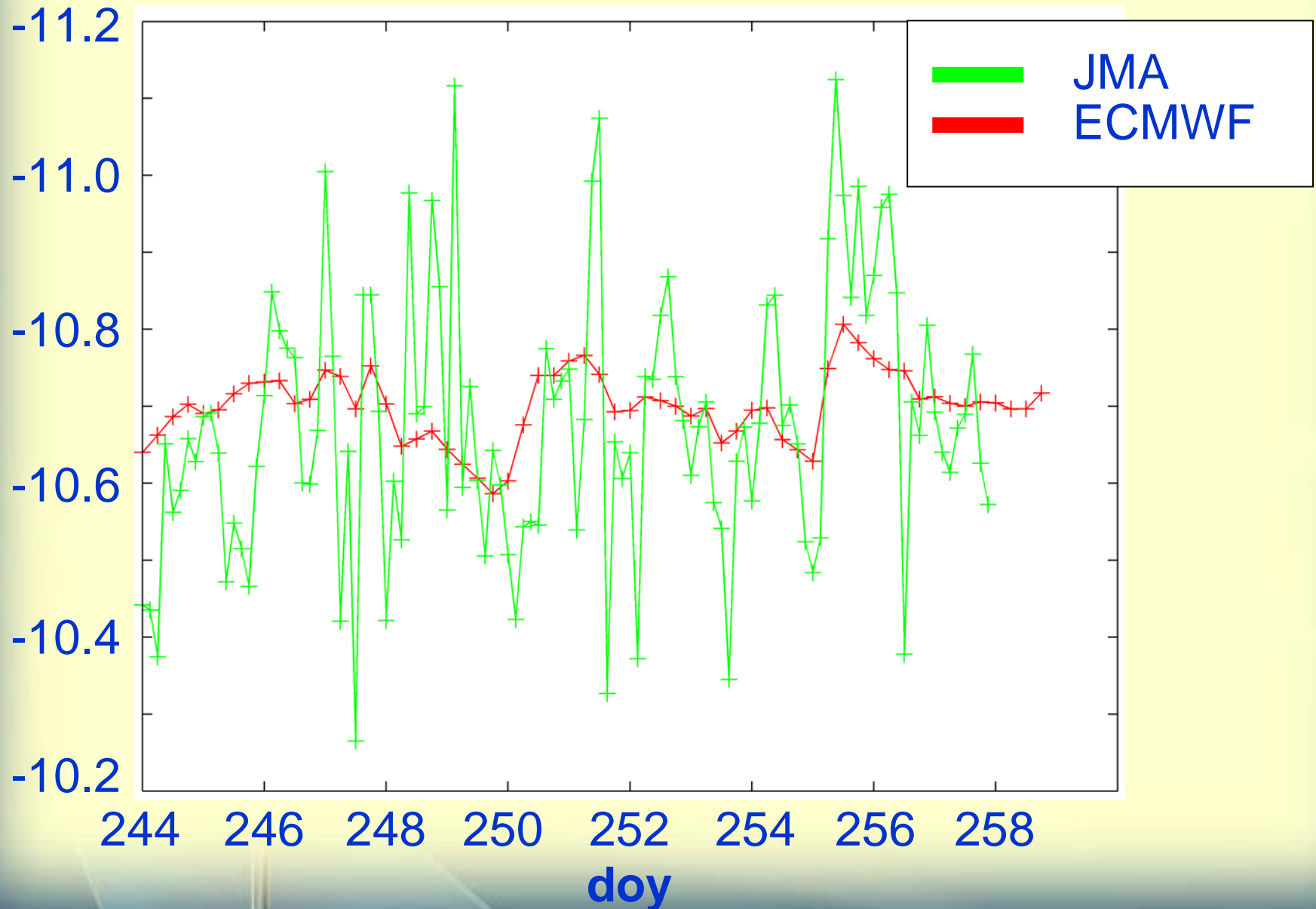
mean of a at 3°



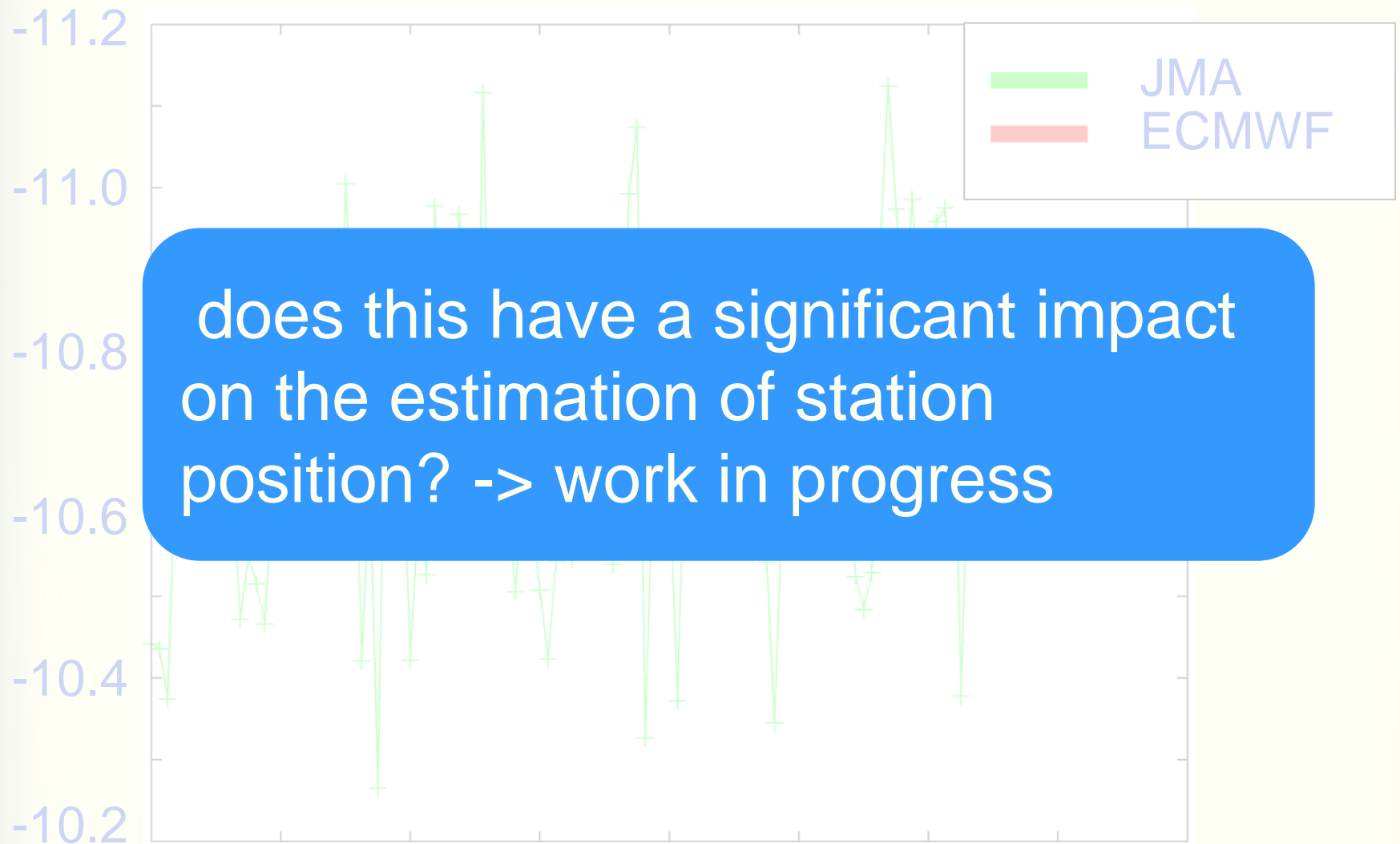
a from ECMWF (mean)



mapping function @ 5 degree elevation



wet delay @ 5 degree elevation

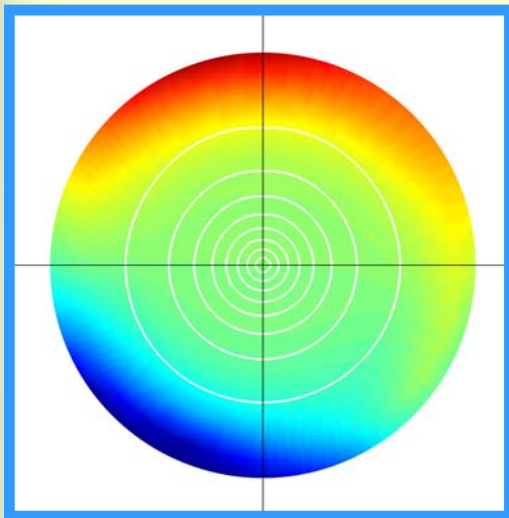


Modeling azimuthal asymmetries

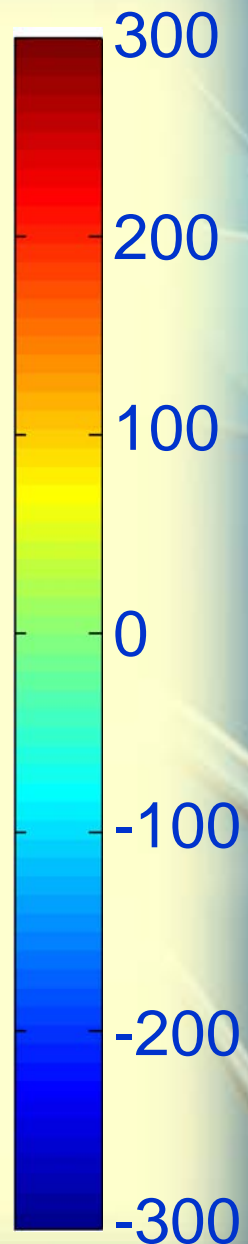
„classical“ gradients (IERS Conventions 2003)

$$\Delta L_{\text{asymm}} = m f_g(e) \cdot [G_N \cdot \cos(\alpha) + G_E \cdot \sin(\alpha)]$$

day 245, 06:00 UT [mm]



without modeling
asymmetries

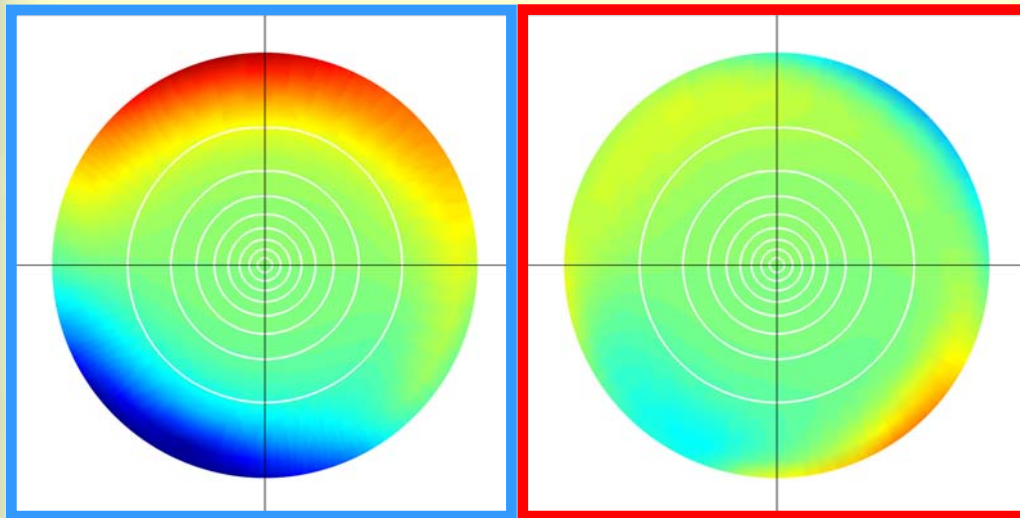


Modeling azimuthal asymmetries

„classical“ gradients (IERS Conventions 2003)

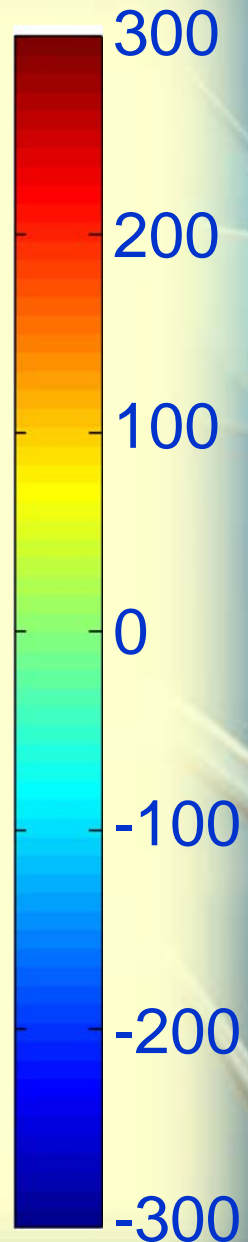
$$\Delta L_{\text{asymm}} = m f_g(e) \cdot [G_N \cdot \cos(\alpha) + G_E \cdot \sin(\alpha)]$$

day 245, 06:00 UT [mm]



without modeling
asymmetries

classical gradients
applied

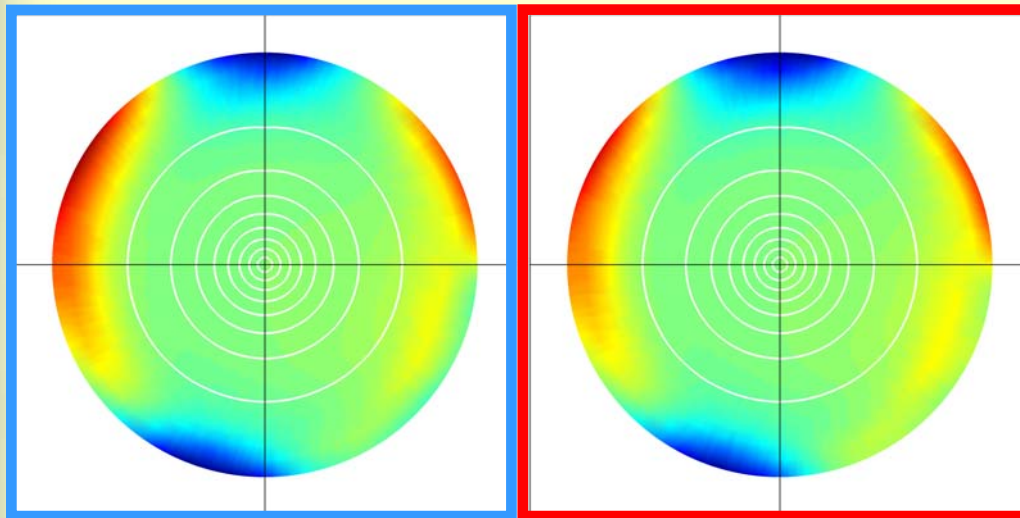


Modeling azimuthal asymmetries

„classical“ gradients (IERS Conventions 2003)

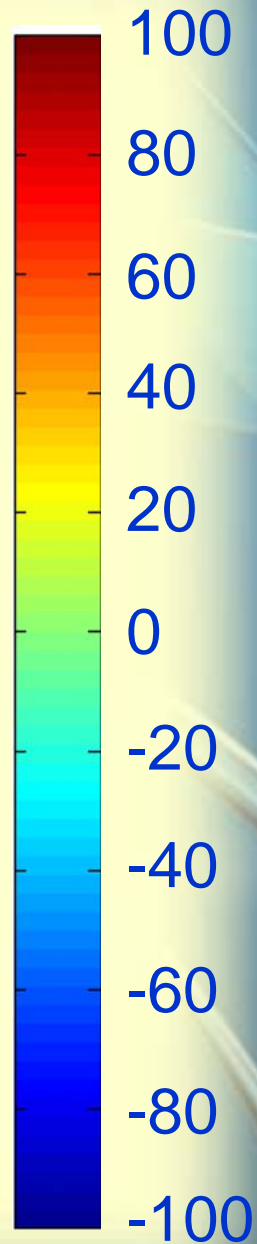
$$\Delta L_{\text{asymm}} = m f_g(e) \cdot [G_N \cdot \cos(\alpha) + G_E \cdot \sin(\alpha)]$$

day 244, 09:00 UT [mm]



without modeling
asymmetries

classical gradients
applied

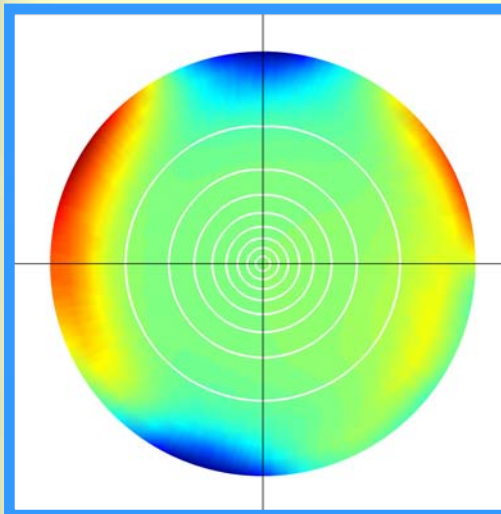


Modeling azimuthal asymmetries

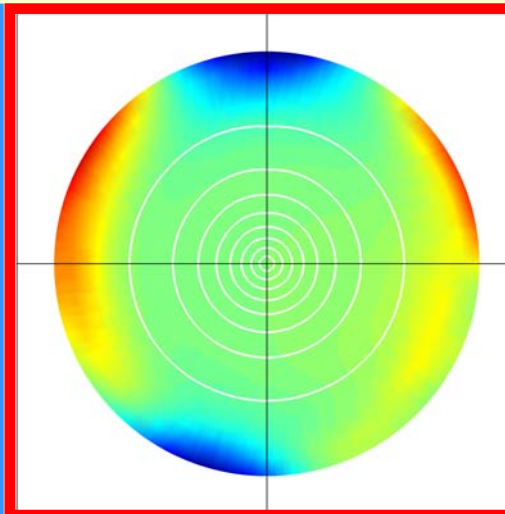
„double“ gradients

$$\Delta L_{\text{asymm}} = m f_g(e) \cdot [G_N \cdot \cos(2\alpha) + G_E \cdot \sin(2\alpha)]$$

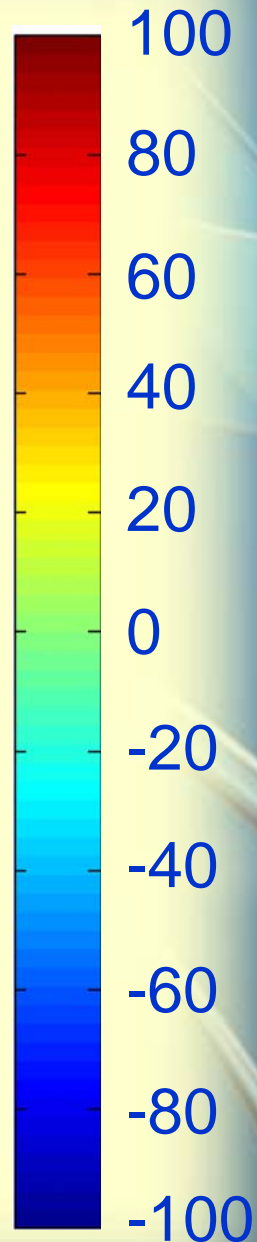
day 244, 09:00 UT [mm]



without modeling
asymmetries



classical gradients
applied

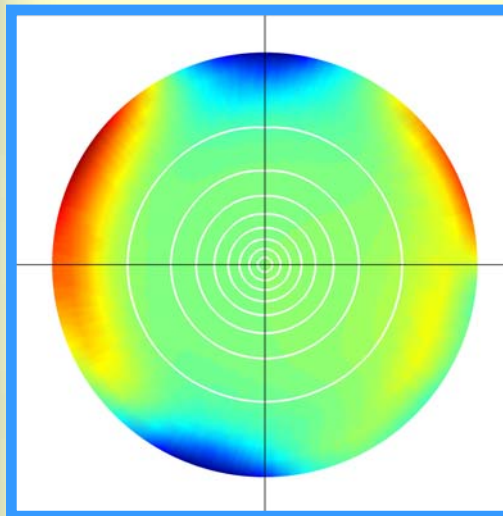


Modeling azimuthal asymmetries

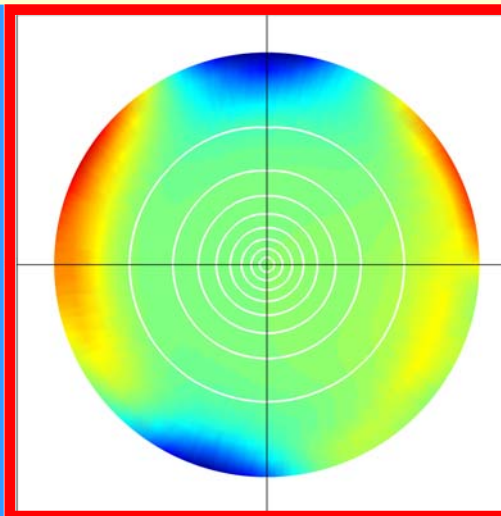
„double“ gradients

$$\Delta L_{\text{asymm}} = m f_g(e) \cdot [G_N \cdot \cos(2\alpha) + G_E \cdot \sin(2\alpha)]$$

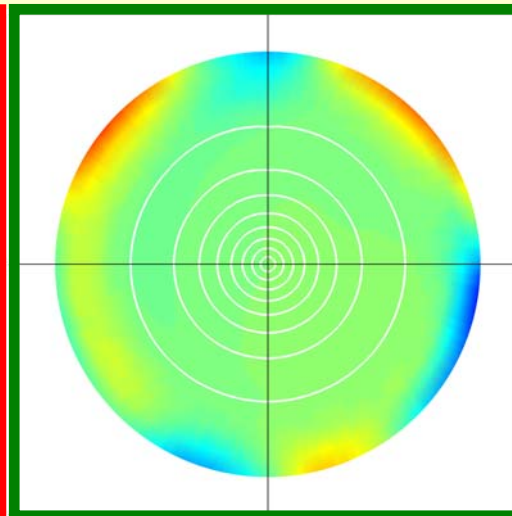
day 244, 09:00 UT [mm]



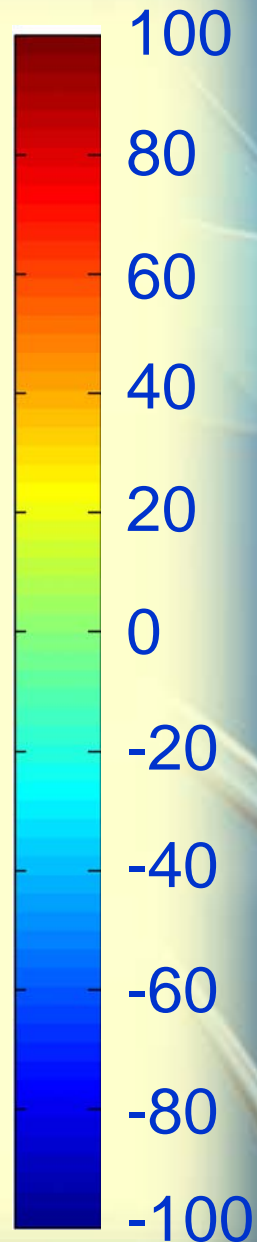
without modeling
asymmetries



classical gradients
applied



double gradients
applied



Summary

- ***a***-coefficient derived from KARAT delays shows significantly more variation than the ***a***-coefficient determined from ECMWF (-> impact?)

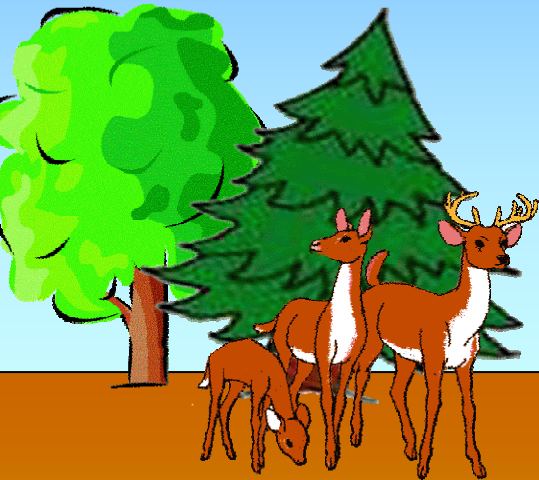
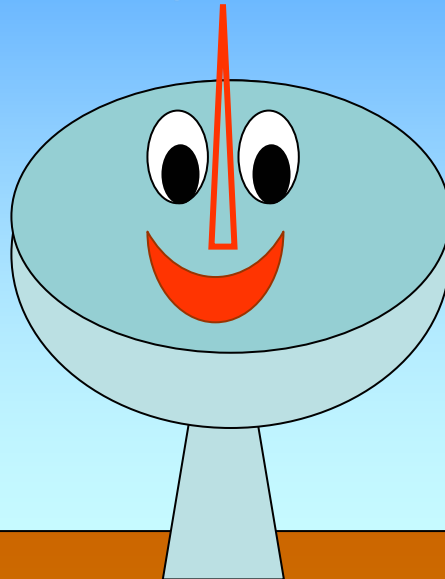
Summary

- a -coefficient derived from KARAT delays shows significantly more variation than the a -coefficient determined from ECMWF (-> impact?)
- the asymmetric delay is dependent on the method of determining the mapping function coefficients

Summary

- α -coefficient derived from KARAT delays shows significantly more variation than the α -coefficient determined from ECMWF (-> impact?)
- the asymmetric delay is dependent on the method of determining the mapping function coefficients
- asymmetric delays exhibit not only gradient like behaviour -> possibilities of a better modeling have to be explored in more detail

Thanks for your attention!



The author is recipient of a DOC-fORTE fellowship of the Austrian Academy of Sciences at the Institute of Geodesy and Geophysics, Vienna University of Technology.